

Swelling Porous Media: Developing a Multiscale Model of Overburden Pressure as a Function of Water Content in Montmorillonite-Bearing Clayey Soils

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Outline

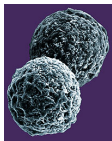
- 1 Swelling Porous Media
- 2 Modeling Approach - HMT
- 3 Macroscale Equations
- 4 Generalized Terzaghi Stress Principle
- 5 Overburden Pressure vs Water Content for Clayey Soils

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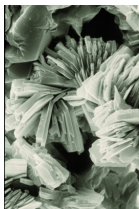
Examples of Swelling Porous Media

Swelling Porous Media: Examples



Drug delivery polymers

R. Langer, Massachusetts Institute of Technology, Cambridge, MA



Clay



Rye Bread

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Obtaining Macroscale Governing Equations

Homogenization or Averaging (Auriault, Whitaker):

Microscale	Avg/Homogenize	Macroscale
Balance Laws	----->	Balance Laws
Const. Eqns		Const. Eqns

Mixture Theory (Bowen):

Postulate Balance	Exploit	Macroscale
Laws at	Entropy Ineq.	Const. Eqns
Macroscale	----->	

Hybrid Mixture Theory (Hassanizadeh and Gray):

Microscale	Volume Average	Macroscale
Balance Laws	----->	Balance Laws

Exploit		Macroscale
Entropy Ineq.		Const. Eqns
----->		

Exploiting Entropy Inequality for a Fluid

Conservation of Linear Momentum

$$\frac{\partial}{\partial t}(\rho \mathbf{v}) + \nabla \cdot (\rho \mathbf{v} \mathbf{v}) - \nabla \cdot \mathbf{t} - \rho \mathbf{g} = 0$$

Need constitutive equation for stress, \mathbf{t} .

Assume a set of independent variables upon which all constitutive variables depend:

$$T, \rho, (\nabla \mathbf{v})_{\text{sym}} = \mathbf{d}$$

Exploiting entropy inequality gives

Exploiting Entropy Inequality

Exploiting entropy inequality, linearizing about equilibrium ($\nabla \mathbf{v} = \mathbf{0}$):

$$\mathbf{t} = (-\bar{p} + \lambda \text{tr}(\mathbf{d}))\mathbf{I} + 2\nu : \mathbf{d}, \quad \bar{p} = \rho^2 \frac{\partial \Psi}{\partial \rho}$$

where Ψ is the Helmholtz potential.

It is generally *assumed* that

$$-\frac{1}{3}\text{tr}(\mathbf{t}) = p = \bar{p} = \rho^2 \frac{\partial \Psi}{\partial \rho}$$

Substitute in Conservation of Linear Momentum: **N-S Eqns:**

$$\rho \frac{D\mathbf{v}}{Dt} + \nabla p - \nabla(\lambda \nabla \cdot \mathbf{v}) - \nabla \cdot \left(\mu \frac{1}{2} (\nabla \mathbf{v} + (\nabla \mathbf{v})^T) \right) - \rho \mathbf{g} = \mathbf{0}$$

It's often assumed that $\lambda = -\frac{2}{3}\mu$

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Macroscale Balance Laws

$$\text{Volume Fraction} = \varepsilon^\alpha = \frac{|\delta V_\alpha|}{|\delta V|}, \quad \varepsilon^l + \varepsilon^s = 1$$

Conservation of Mass ($\alpha = l, s$)

$$\frac{D^\alpha(\varepsilon^\alpha \rho^\alpha)}{Dt} + \varepsilon^\alpha \rho^\alpha \nabla \cdot \mathbf{v}^\alpha = \sum_{\beta \neq \alpha} \hat{\mathbf{e}}_\beta^\alpha$$

Conservation of Momentum ($\alpha = l, s$)

$$\varepsilon^\alpha \rho^\alpha \frac{D^\alpha \mathbf{v}^\alpha}{Dt} - \nabla \cdot (\varepsilon^\alpha \mathbf{t}^\alpha) - \varepsilon^\alpha \rho^\alpha \mathbf{g}^\alpha = \sum_{\beta \neq \alpha} \hat{\mathbf{T}}_\beta^\alpha$$

Conservation of energy ($\alpha = l, s$)
(similarly)

Assumptions on Behavior of Material

Assume:

- Two phases - liquid and solid
- Phases may be charged
- Multiple constituents for each phase

It is assumed that the constitutive variables depend on:

$$\varepsilon^l, \rho^j, \rho^{sj}, T, \mathbf{E}^s, \mathbf{v}^{l,s}, \mathbf{u}^j, \mathcal{E}, \\ \nabla \varepsilon^l, \nabla \rho^{\alpha j}, \nabla T, \nabla \mathbf{E}^s, \mathbf{d}^l, \nabla \mathbf{u}^j,$$

ε^l = void fraction

ρ^j = density of constituent j in liquid phase
(mass l_j / volume of liquid)

T = temperature

$\mathbf{v}^{l,s}$ = relative velocities: $\mathbf{v}^l - \mathbf{v}^s$

\mathbf{u}^j = diffusive velocity: $\mathbf{v}^j - \mathbf{v}^l$

\mathcal{E} = electric field

\mathbf{E}^s = macroscale strain of solid phase

Closure issue: ε^l is a constitutive variable.

Three Pressures

As a result of having volume fraction, ε^l , as an independent variable:

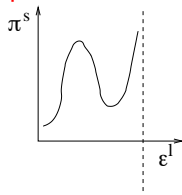
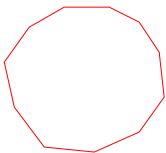
$$\begin{aligned} p^l &= -\frac{1}{3}\text{tr}(\mathbf{t}^l) & \bar{p}^l &= (\rho^l)^2 \left. \frac{\partial \Psi^l}{\partial \rho^l} \right|_{\varepsilon^l} & \pi^l &= \varepsilon^l \rho^l \left. \frac{\partial \Psi^l}{\partial \varepsilon^l} \right|_{\rho^l} \\ p^s &= -\frac{1}{3}\text{tr}(\mathbf{t}^s) & \bar{p}^s &= (\rho^s)^2 \left. \frac{\partial \Psi^s}{\partial \rho^s} \right|_{\varepsilon^s} & \pi^s &= \varepsilon^s \rho^s \left. \frac{\partial \Psi^s}{\partial \varepsilon^s} \right|_{\rho^s} \end{aligned}$$

where we can show mathematically (no assumptions)

$$p^\alpha = \bar{p}^\alpha + \pi^\alpha, \quad \alpha = l, s$$

Three Pressures

What is $\pi^s = \varepsilon^s \rho^s \left. \frac{\partial \Psi^s}{\partial \varepsilon^s} \right|_{\rho^s, \rho^l}$? Represents how the solid energy changes when changing porosity without changing the densities of liquid or solid. **Configurational pressure**



Three Pressures

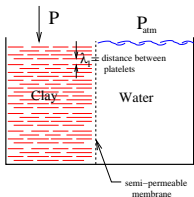
What is $\pi^l = \varepsilon^l \rho^l \left. \frac{\partial \Psi^l}{\partial \varepsilon^l} \right|_{\rho^s, \rho^l}$? Represents how the liquid energy changes as the volume fraction changes

Non-swelling porous media, $\pi^l = 0$: we recover $p^l = \bar{p}^l$.

Swelling porous media, π^l can be measured using a reverse osmotic experiment:

$$P = \pi^l = \exp\left(\frac{\alpha}{\lambda}\right) - 1$$

α is a function of cation exchange and specific surface area



P.F. Low, *Soil Science Society of America Journal*, 44(4), 1980

P.F. Low, *Soil Science Society of America Journal*, 45, 1981

P.F. Low, *Progr Colloid & Polym Sci*, 95, 1994

Three Pressures

Consequences

- Low's empirical results can be derived by balancing the *macroscale* Gibbs (chemical) potential across the boundary.
- **The boundary condition between vicinal water and bulk water is *not* a balance of pressures. It is a balance of chemical potentials.**
- DLVO theory was originally derived for gas particles and has been applied to vicinal liquid in swelling clay at the *microscale*. According to Mitchel and Soga (2005), a double layer model cannot “explain the swelling of pure clay accurately.”

Darcy's Law

Conservation of Linear Momentum for liquid phase

$$\varepsilon^l \rho^l \frac{D^l \mathbf{v}^l}{Dt} - \nabla \cdot (\varepsilon^l \mathbf{t}^l) - \varepsilon^l \rho^l \mathbf{g} = \sum_{\beta \neq \alpha} \hat{\mathbf{T}}_s^l$$

Constitutive results from exploiting entropy inequality (linearizing)

$$\mathbf{t}^l = -p^l + \mu \mathbf{d}^l$$

$$\hat{\mathbf{T}}_s^l = p^l \nabla \varepsilon^l - (\varepsilon^l)^2 \mathbf{R} \cdot \mathbf{v}^{l,s}$$

Neglect gravity, inertial term, and viscous term:

$$\mathbf{R} \cdot \mathbf{v}^{l,s} = -\varepsilon^l \nabla p^l - \pi^l \nabla \varepsilon^l$$

Flow can be inhibited by swelling.

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What is Terzaghi Stress Principle with π^l and π^s ?

By adding the conservation of momentum we have that the total or applied stress, \mathbf{t} ,

$$\mathbf{t} = \varepsilon^l \mathbf{t}^l + \varepsilon^s \mathbf{t}^s$$

$$\rho = \varepsilon^l \rho^l + \varepsilon^s \rho^s \quad \text{where} \quad \rho^l = -\frac{1}{3} \text{tr} \mathbf{t}^l, \quad \rho^s = -\frac{1}{3} \text{tr} \mathbf{t}^s$$

Experimentally we have (for nonswelling porous media)

Terzaghi Stress Principle:

$$\mathbf{t}_{\text{eff}} = \mathbf{t} + \alpha \rho^l \mathbf{I} \quad \rho_{\text{eff}} = \rho - \alpha \rho^l$$

When $\alpha = 1$ it can be shown that the max. microscale shear stress is proportional to the Terzaghi effective pressure, ρ_{eff}

(Schreyer et al., 2021)

What is Terzaghi Stress Principle with π^l and π^s ?

Using $\bar{p}^l = \bar{p}^s$ at equilibrium (result from exploiting entropy inequality) we get

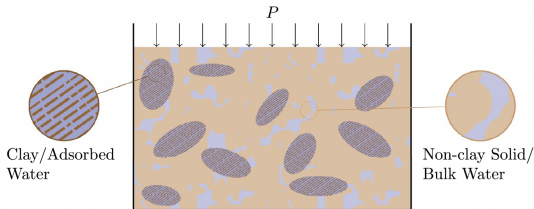
$$\begin{aligned}
 p_{\text{eff}} &= p - p^l \\
 &= (\varepsilon^l p^l + \varepsilon^s p^s) - (\varepsilon^l + \varepsilon^s) p^l \\
 &= \varepsilon^l (\bar{p}^l + \pi^l) + \varepsilon^s (\bar{p}^s + \pi^s) - \varepsilon^l (\bar{p}^l + \pi^l) - \varepsilon^s (\bar{p}^l + \pi^l) \\
 &= \varepsilon^s (\pi^s - \pi^l)
 \end{aligned}$$

So this theory says that the **effective pressure is completely determined by the swelling pressure and configurational pressure** (see also, Ehlers 2018 for multiphase flow where saturation is used instead of volume fraction).

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Generalized Terzaghi Stress Principle



For soils containing swelling clays, we want an equation that reduces to Terzaghi's when there is no swelling clays and reduces to Low's relationship when it is all swelling clay.

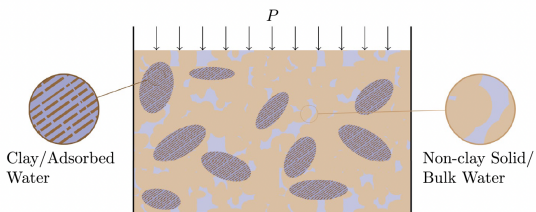
This is a 3-scale problem:

microscale Distinguish between the vicinal (adsorbed) fluid and clay platelets

mesoscale Distinguish between the clay particles and the **bulk fluid**

macroscale Cannot distinguish between the phases.

Generalized Terzaghi Stress Principle



Four phases:

- A_f vicinal fluid
- A_s clay minerals that adsorb fluid
- B_f nonabsorbed fluid
- B_s non-clay solids

Generalized Terzaghi Stress Principle

Results from HMT:

$$\pi_{\beta}^{\alpha} = \sum_{\delta} \varepsilon^{\delta} \rho^{\delta} \left. \frac{\partial \psi^{\delta}}{\partial \varepsilon^{\alpha}} \right|_{\varepsilon^{\gamma}, \rho^{\zeta}}$$

$$\bar{p}_{\beta}^{\alpha} = - \sum_{\delta} \varepsilon^{\delta} \rho^{\delta} \left. \frac{\partial \psi^{\delta}}{\partial \varepsilon^{\alpha}} \right|_{\varepsilon^{\beta}, \varepsilon^{\zeta}, \rho^{\zeta}}$$

$$\gamma \neq \alpha, \beta, \quad \alpha, \zeta = A_l, A_s, B_l, B_s$$

$$p^{\alpha} = -\frac{1}{3} \text{tr} \mathbf{t}^{\alpha} = (\rho^{\alpha})^2 \frac{\partial \psi^{\alpha}}{\partial \rho^{\alpha}}$$

In words: $\pi_{B_l}^{A_l}$ is the change in energy as water moves between A_l and B_l while volume fraction of solid phases are fixed and all densities stay fixed.

π_{β}^{α} is the swelling pressure above the bulk phase pressure so we can use Low's exponential relationship (almost).

Generalized Terzaghi Stress Principle

Results from HMT: Equivalent 3 pressure relationships:

$$p^\alpha - p^\beta = \pi_\beta^\alpha + \bar{p}_\beta^\alpha$$

$$\sum_\beta \varepsilon^\beta \rho^\beta \frac{\partial \psi^\beta}{\partial \varepsilon^\alpha} - p^\alpha + p^{B_s} = 0 \quad \Rightarrow$$

$$\bar{p}_{B_s}^\alpha = 0 \quad \text{at equil}$$

And using the fact that the total pressure is the weighted sum of the 4 pressures:

$$p = \varepsilon^{A_l} p^{A_l} + \varepsilon^{A_s} p^{A_s} + \varepsilon^{B_l} p^{B_l} + \varepsilon^{B_s} p^{B_s}$$

We get...

Generalized Terzaghi Stress Principle

Effective pressure and total pressure expressions are

$$p_{\text{eff}} = \varepsilon^{A_s} \pi_{A_l}^{A_s} + \varepsilon^{B_s} \pi_{B_l}^{B_s}$$

$$p = (\varepsilon^{A_s} + \varepsilon^{A_l}) \pi_{B_l}^{A_l} + p^{B_l} + p_{\text{eff}}$$

where the terms in blue act as the effective fluid pore pressure.
 How do we use this generalized Terzaghi Stress Principle? This is the art of modeling! Make some educated assumptions!

Generalized Terzaghi Stress Principle

Assume

- 1 p^{B_l} , bulk fluid pressure, and p , overburden pressure, mass fractions of A_s and B_s , and total water content are measurable
- 2 adsorbed and bulk water have the same density, $\rho^{A_l} = \rho^{B_l}$ (see Low for the magnitude of this approximation)
- 3 Low's exponential relationship is still valid for $\pi_{B_l}^{A_l}$
- 4 effective pressure, p_{eff} is linear in strain
- 5 solid components are incompressible, $V^{A_s} = V_0^{A_s}$ and $V^{B_s} = V_0^{B_s}$

After a fair bit of algebra:

$$p = P_{\text{atm}} \frac{\omega^A + f^c \frac{\rho^l}{\rho^{A_s}}}{\omega + \frac{\rho^l}{\rho^s}} \left(\exp \left(\frac{\alpha f^c}{\omega^A} \right) - 1 \right) + p^{B_l} - K \frac{\omega - \omega_0}{\omega_0 + \frac{\rho^l}{\rho^s}}$$

Generalized Terzaghi Stress Principle

$$p = P_{\text{atm}} \frac{\omega^A + f^c \frac{\rho^l}{\rho^s}}{\omega + \frac{\rho^l}{\rho^s}} \left(\exp \left(\frac{\alpha f^c}{\omega^A} \right) - 1 \right) + p^{B_l} - K \frac{\omega - \omega_0}{\omega_0 + \frac{\rho^l}{\rho^s}}$$

= **Low's relationship** + **bulk liquid pressure** (= p_{atm} under drained conditions) + **bulk linear elasticity** where

ω^A ratio of mass of adsorbed water to mass of total solid

ω ratio of mass of total water to mass of total solid

f^c ratio of mass of clay to mass of total solid

K bulk modulus of elasticity for effective pressure, assumed constant

α is from Low's exponential relationship and is a fn of cation exchange capacity and specific surface area

Generalized Terzaghi Stress Principle

What to do with adsorbed (vicinal) water content?

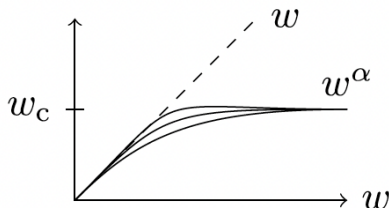


Figure: Adsorbed water content, $w^\alpha = w^A$, as a function of total water

For results, we assumed piecewise linear

Generalized Terzaghi Stress Principle

We now have an equation with 4 empirical constants:

f^c mass fraction of clay

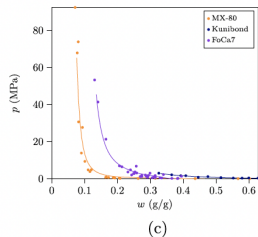
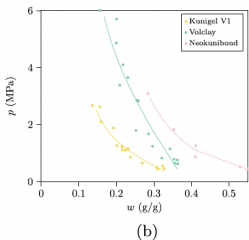
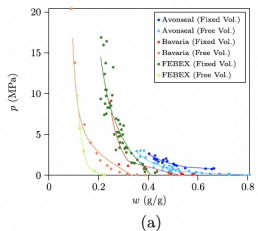
ω_c from ω vs ω^A function

ω_0 initial water content,

K bulk modulus of elasticity for effective pressure

$$p = P_{\text{atm}} \frac{\omega^A(\omega, \omega_c) + f^c \frac{\rho^l}{\rho^s}}{\omega + \frac{\rho^l}{\rho^s}} \left(\exp \left(\frac{\alpha \pi f^c}{4 \omega^A(\omega, \omega_c)} \right) - 1 \right) + p^{B_l} - K \frac{\omega - \omega_0}{\omega_0 + \frac{\rho^l}{\rho^s}}$$

Generalized Terzaghi Stress Principle for Clayey soils - results



Conclusions

- When modeling swelling porous media, **swelling pressure is a key process to incorporate.**
- The proposed model is designed to **capture the behavior of soils ranging from low pressure (dominated by skeletal deformation), to high pressure limits where adsorbed water pressure dominates.**
- The assumption that **the bulk modulus is constant can be relaxed** to be piecewise constant as is done for modeling plasticity [e.g. Schreyer et al., 2021]
- **Comparison with machine learning models:** coefficients have physical interpretation, allowing for predictions for clayey solid not used in the training set.
- **Other physics models have fewer parameters**, but are designed for the low moisture content regime only.

THANK YOU

References

W. Ehlers, *Effective stresses in multiphase porous media: a thermodynamic investigation of a fully non-linear model with compressible and incompressible constituents*, Geomechanics for Energy and the Environment, 15, 2018

Hassanizadeh M, Gray WG *General conservation equations for multi-phase systems: 1. Averaging procedure* Adv Water Resour 2:131?144. [https://doi.org/10.1016/0309-1708\(79\)90025-3](https://doi.org/10.1016/0309-1708(79)90025-3), 1979

Hassanizadeh M, Gray WG *General conservation equations for multi-phase systems: 2. Mass, momenta, energy, and entropy equations* Adv Water Resour 2:191?203. [https://doi.org/10.1016/0309-1708\(79\)90035-6](https://doi.org/10.1016/0309-1708(79)90035-6), 1979

Hassanizadeh M, Gray WG *General conservation equations for multi-phase systems: 3. Constitutive theory for porous media flow* Adv Water Resour 3(1):25-40, 1980

Low PF (1980) *The swelling of clay: II. Montmorillonites*, Soil Sci Soc Am J 44(4):667?676. <https://doi.org/10.2136/sssaj1980.03615995004400040001x>, 1980

Low PF *The swelling of clay: III. Dissociation of exchangeable cations*, Soil Sci Soc Am J 45(6):1074? 1078. <https://doi.org/10.2136/sssaj1981.03615995004500060013x>, 1981

Low PF, Margheim JF *The swelling of clay: I. Basic concepts and empirical equations*, Soil Sci Soc Am J 43(3):473?481. <https://doi.org/10.2136/sssaj1979.03615995004300030010x>, 1979

Mašin D, Khalili N *Swelling phenomena and effective stress in compacted expansive clays*, Can Geotech J 53(1):134?147. <https://doi.org/10.1139/cgj-2014-0479>, 2016

Mitchell JK, Soga K *Fundamentals of soil behavior*, 3rd edn. Wiley, New York, 2005

H.L. Schreyer, B.C. Lampe, L.G. Schreyer, J.C. Stormont *Microscale Analysis Demonstrating the Significance of Shear and Porosity in Hydrostatic Compression of Porous Media*, International Journal of Rock Mechanics and Mining Sciences, 145, 104751, <https://doi.org/10.1016/j.ijrmms.2021.104751>, 2021

R. Whitehead, L. Schreyer, I. Akin *A Multi-scale Model of Overburden Pressure and Water Content in Montmorillonite-bearing Clayey Soils*, Geotechnical and Geological Engineering, 42, 3843-3856 doi.org/10.1007/s10706-024-02761-0, 2024