

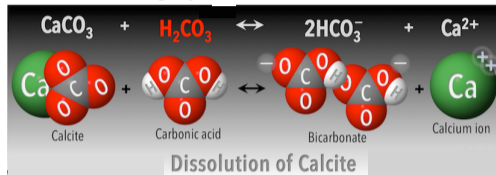
Model for Flow in Karstified Carbonate Reservoirs using Two-Level Upscaling and Surrogate-Based Transmissibilities under Geomechanical Effects

Patrícia Pereira, Tayná Lobo,
Josué Barroso, Emanuel Gomes,
Márcio Murad



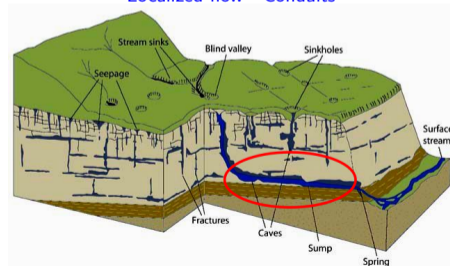
1. Motivation
2. Literature Review
3. MsEDKM
4. MsEDKM coupled with Geomechanics
5. Computational Results
6. Wrap-up

- Sedimentary carbonate rocks are highly reactive to fluids



- Dissolution processes \implies karst:
vugs, enlarged fractures, caves, conduit networks
- Karst in carbonate reservoirs:
 - impact reservoir storage capacity and subsurface fluid flow
 - collapse of wells
 - fluid loss during drilling
- Models need to incorporate these complexities

Localized flow - Conduits



Runkel, A. C. et al, (2003).

- 2 ϕ 2K with enhanced permeabilities (Machado, 2020)
- Coupled 3D-1D Discrete Model (Murad, 2020; Ferraz,2021; Landim,2023,2025)
- Triple Porosity Models (Kang, 2004)
- EDKM (Correia at al, 2018)
 - Extension of the well established EDFM (Li and Lee, 2008; Moinfar et al. 2014)
 - Single porosity model with two grid domains (matrix and karst properties)
 - Communication between matrix/karst via special connections (NNC)
 - **Karst representation: same block size as refined geological grid**
 - **Transmissibility matrix/karst computed analytically.**

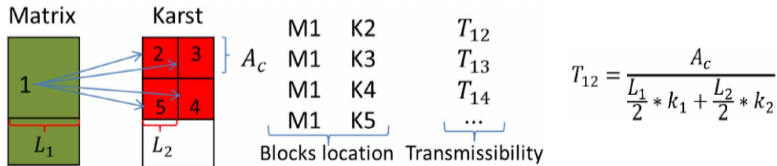
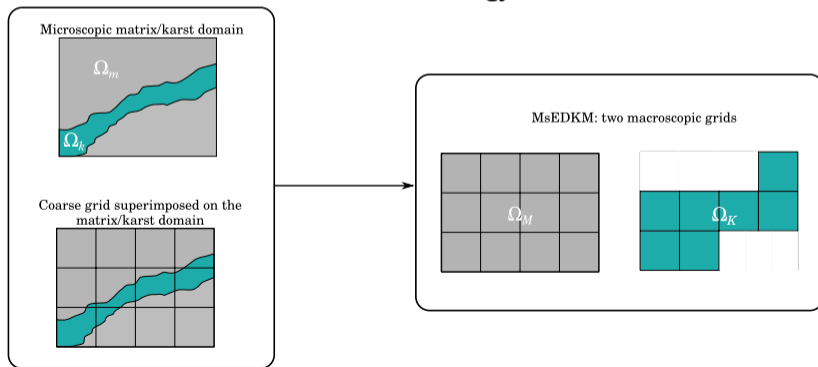
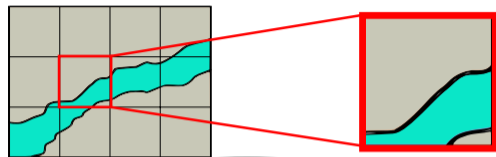
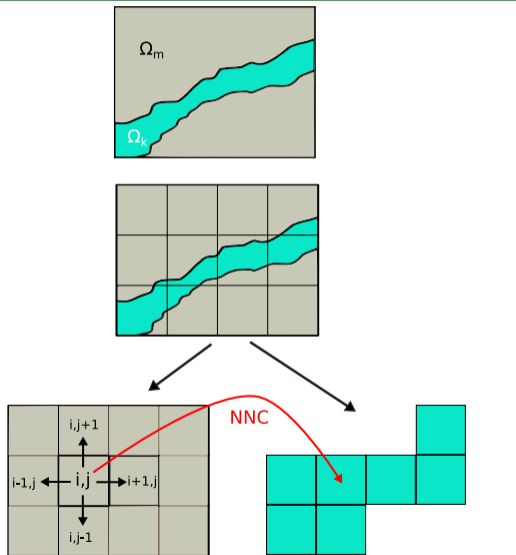


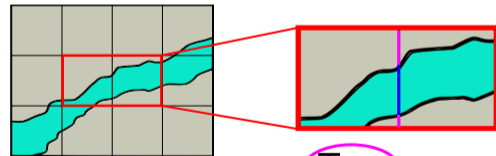
Figure 3—Application of special connections.

- MsEDKM: multiscale embedded discrete karst model
 - **Karst representation: same block size as the matrix grid**
 - **Matrix-karst transmissibilities computed via numerical simulations**
 - **Surrogate model trained on high-fidelity numerical simulations to reduce computational cost and enable scalable use of the methodology**





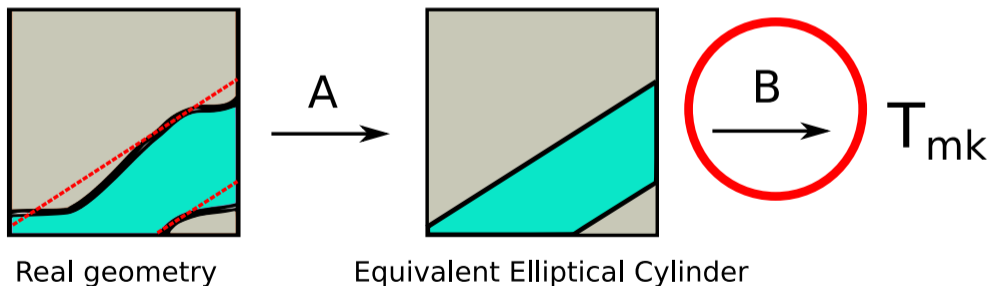
mass exchange between matrix and karst: T_{mk}



T_{mm}

T_{kk}

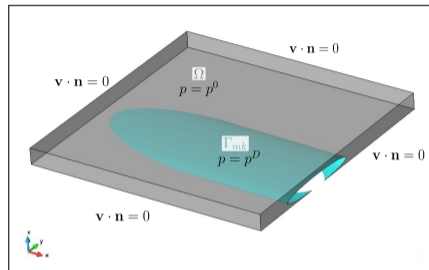
- **A.** Moment of inertia based upscaling (micro-meso)
Goal: regularization of the conduit surface roughness
- **B.** Flow based upscaling (meso-macro)
Goal: computation of hydraulic transmissibility between karst and matrix (T_{mk})



Numerical Computation of T_{mk} ($K_k \gg K_m$): No Geomechanical Effects

Solve the local microscopic transient problem:

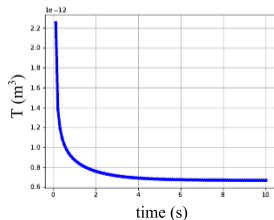
$$\left\{ \begin{array}{ll} \beta \frac{\partial p}{\partial t} + \nabla \cdot \mathbf{v} = 0 & \text{in } \Omega_m \\ \mathbf{v} = -\frac{\mathbf{K}}{\mu} \nabla p & \text{in } \Omega_m \\ p(\mathbf{x}, 0) = p^0(\mathbf{x}) & \text{in } \Omega_m \\ p(\mathbf{x}, t) = p^D(\mathbf{x}) & \text{on } \Gamma_{mk} \\ \mathbf{v} \cdot \mathbf{n} = 0 & \text{on } \Gamma_m \end{array} \right.$$



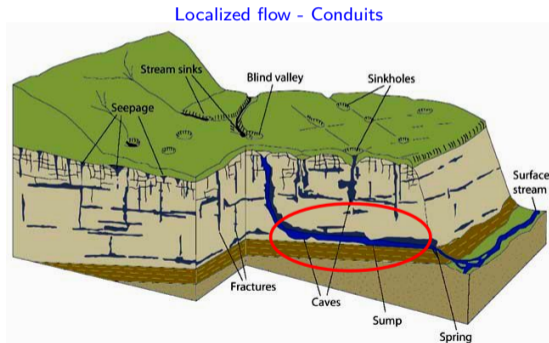
Compute T_{mk} :

$$T_{mk}(t) = \frac{\mu q}{\bar{p} - p^D}$$

$$T_{mk} = \lim_{t \rightarrow \infty} T_{mk}(t)$$



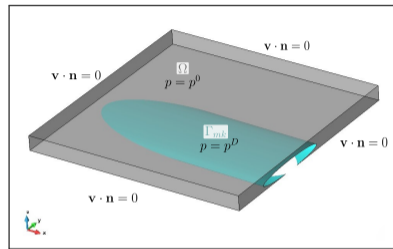
Extend the MsEDKM framework to account for geomechanical effects on flow through stress-dependent hydraulic properties.



Runkel, A. C. et al, (2003).

Solve the **non-linear** local microscopic transient problem:

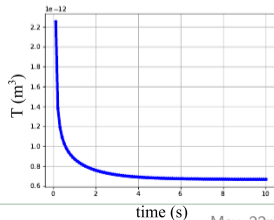
$$\left\{ \begin{array}{ll} \beta_{\text{eff}} \frac{\partial p}{\partial t} + \nabla \cdot \mathbf{v} = 0 & \text{in } \Omega_m \\ \mathbf{v} = -\frac{\mathbf{K}}{\mu} \nabla p & \text{in } \Omega_m \\ \mathbf{K} = \mathbf{K}(\phi(p)) & \\ p(\mathbf{x}, 0) = p^0(\mathbf{x}) & \in \Omega_m \\ p(\mathbf{x}, t) = p^D(\mathbf{x}) & \text{on } \Gamma_{mk} \\ \mathbf{v} \cdot \mathbf{n} = 0 & \text{on } \Gamma_m \end{array} \right.$$



Compute T_{mk} :

$$T_{mk}(t) = \frac{q\mu}{\bar{p} - p^D}$$

$$T_{mk} = \lim_{t \rightarrow \infty} T_{mk}(t)$$

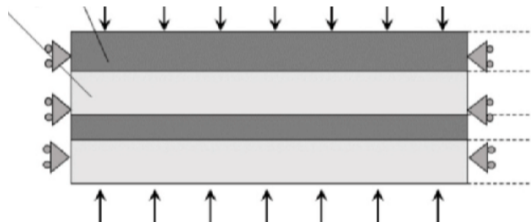


Geomechanical Coupling (analytical approach)

Prescribed Stress Path

Assumptions:

- *In-situ* state (input data): $\{p^{\text{ref}}, \sigma^{\text{ref}}, \mathbf{K}^{\text{ref}}, \phi^{\text{ref}}\}$.
- Total stress main directions are aligned to the coordinate system $\{x, y, z\}$
- **Prescribed total stress path (oedometric)**
- Lateral strains vanish on outer interface
 $\Delta\epsilon_x = 0$
- Constant vertical total stress $\Delta\sigma_V = 0$
- Effective stress changes with pressure
 - $\Delta\sigma'_V = \alpha\Delta p$
 - $\Delta\sigma'_H = \Delta\sigma'_h = \frac{\nu}{1-\nu}\Delta\sigma'_V$
- We are not solving full geomechanics at each step.
Instead, we prescribe how effective stress evolves as pressure changes.



$$\mathbf{K} = \mathbf{K}(p)$$

$$p \longrightarrow \Delta\sigma' \longrightarrow \varepsilon_v \longrightarrow \phi \longrightarrow \mathbf{K}$$

Mechanical response

Considering the oedometric stress path:

$$\Delta\sigma' = \alpha (p - p^{\text{ref}}) \begin{pmatrix} \frac{\nu}{1-\nu} & 0 & 0 \\ 0 & \frac{\nu}{1-\nu} & 0 \\ 0 & 0 & 1 \end{pmatrix}$$

Volumetric Strain:

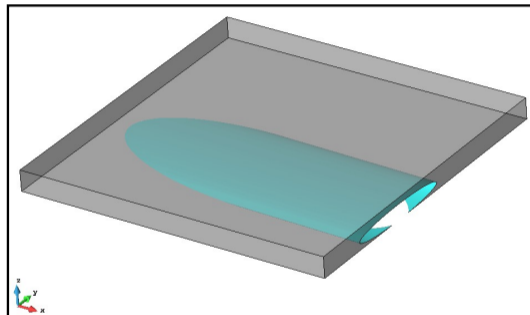
$$\varepsilon_v = \frac{\text{tr}(\Delta\sigma')}{3\mathbb{K}}$$

Porosity and permeability

$$\phi = 1 - (1 - \phi^{\text{ref}}) \exp(-\varepsilon_v)$$

$$\mathbf{K} = \mathbf{K}^{\text{ref}} \left(\frac{\phi}{\phi^{\text{ref}}} \right)^A$$

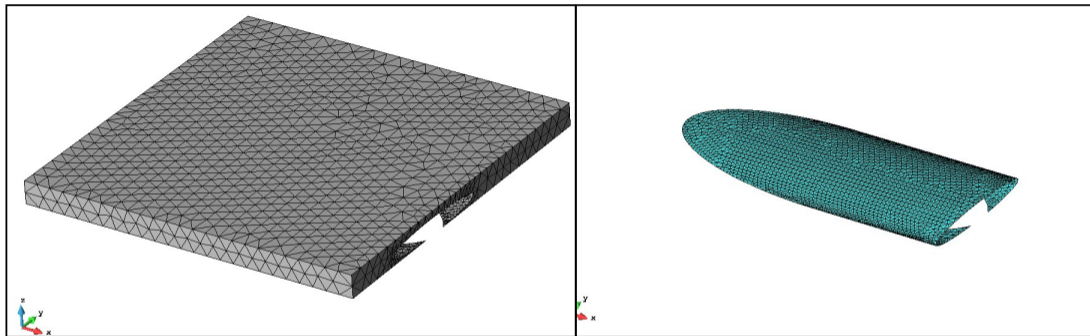
A : permeability-porosity sensitivity exponent.



- cell dimension: 150x150x10 m
- conduit volume fraction: $\approx 10\%$

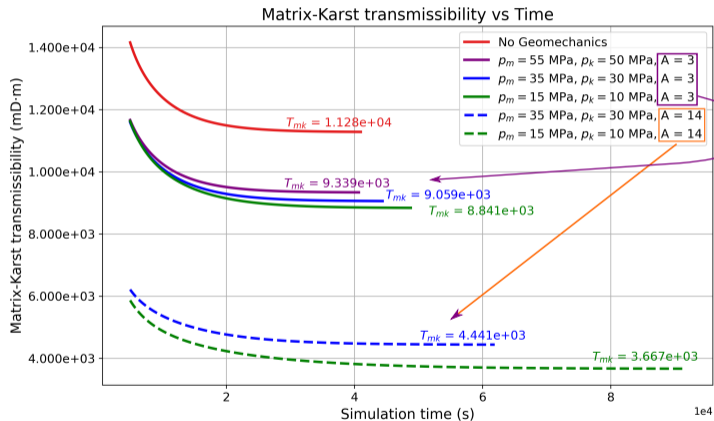
Parameter	In-situ state:	
	Value	Unit (SI)
p^{ref}	60	MPa
σ^{ref}	(80, 80, 100)	MPa
\mathbf{K}^{ref}	(100, 100, 100)	mD
ϕ_{ref}	0.2	–

Input Parameters:		
Parameter	Value	Unit (SI)
E	2.2222×10^9	Pa
ν	0.3	–
\mathbb{K}_s	1.0×10^{32}	Pa
\mathbb{K}_f	2.0×10^9	Pa
α	1.0	–
μ_f	1.0×10^{-3}	Pa s
Δt	1.0×10^2	s

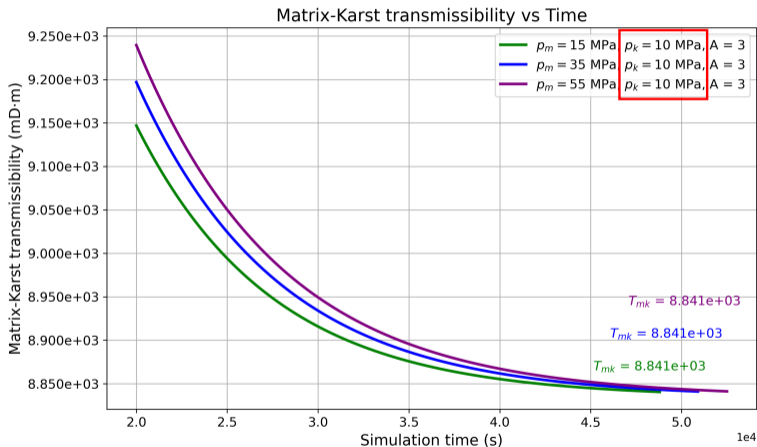


- Finer mesh along the conduit surface, where the flow rate is computed

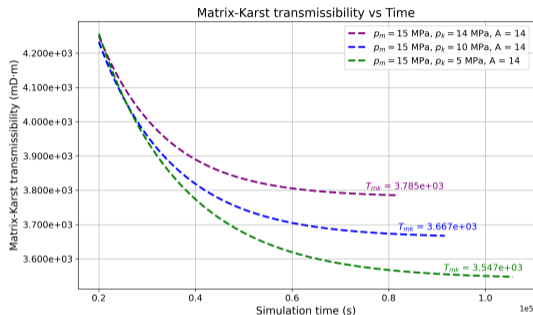
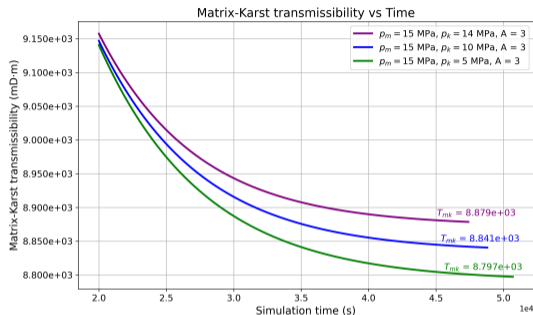
$$T_{mk}(t) = \frac{q\mu}{\bar{p} - p^D}$$



- $\Delta p = (p_m - p_k) = 5 \text{ MPa}$
- A: permeability-porosity sensitivity exponent
- Depletion scenario
- Geomechanical effects reduce transmissibility between matrix/karst transmissibility
- $A \uparrow, \phi \downarrow, \mathbf{K} \downarrow \Rightarrow T_{mk} \downarrow$
- Lower range of pressure: higher geomechanical impact

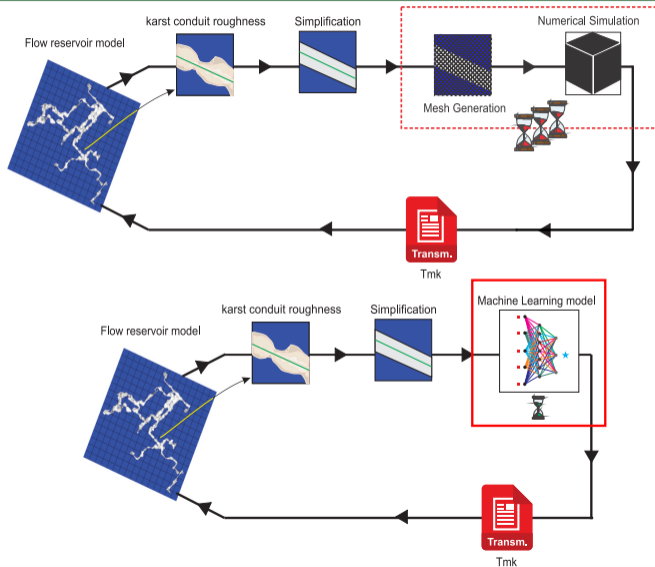


- Depletion scenario
- p_k fixed
- Geomechanical effects:
 $T_{mk} \downarrow$
- T_{mk} are the same for distinct p_m



- p_m fixed
- T_{mk} depends on p_k

$$\Delta \sigma' = \alpha (p - p^{\text{ref}}) \begin{pmatrix} \frac{\nu}{1-\nu} & 0 & 0 \\ 0 & \frac{\nu}{1-\nu} & 0 \\ 0 & 0 & 1 \end{pmatrix}$$



Descriptor vector for the surrogate model

- Geometric and hydraulic attributes

$$\mathbf{X} = \{L_x, L_y, L_z, a_k, b_k, \mathbf{x}_k^{\text{in}}, \mathbf{x}_k^{\text{out}}, L_k, \eta_k, A_{mk}, V_k, \mathbf{K}\}$$

- Geomechanical attributes

$$\mathbf{X}_{\text{geomechanics}} = \{E, \nu, \mathbb{K}_s, \mathbb{K}_f, p^{\text{ref}}, \boldsymbol{\sigma}^{\text{ref}}, \mathbf{K}^{\text{ref}}, A, p_m, p_k\}$$

- Build a new database from hydromechanical numerical simulations
- Construct and train a surrogate model for predicting matrix-karst NNC transmissibilities
- Compute local transmissibilities accounting for geomechanical effects through permeability updates
- Implement the macroscopic flow model incorporating the surrogate model and the local transmissibilities

- MsEDKM: a systematic methodology for integrating karst conduits into reservoir simulators
- Ongoing work extends the MsEDKM framework to account for geomechanical effects through stress-dependent hydraulic properties
- Geomechanical effects are incorporated analytically, under a prescribed stress path, as an alternative to fully coupled hydromechanical numerical simulations
- Matrix–karst mass transfer is computed through numerical simulations, using flow-based upscaling
- ML-based surrogate model will enable large-scale simulations



InterPore2027

Windsor Barra Ho **Rio de Janeiro, Brazil** May 10-13, 2027